

海-陆过渡相与海相页岩气“甜点段”差异特征与形成机理

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摘要: 为了指导海-陆过渡相页岩气的勘探开发, 以鄂尔多斯盆地东缘大宁—吉县地区二叠系山西组2段3亚段(山²₃亚段)海-陆过渡相页岩气甜点段与四川盆地南部地区志留系龙马溪组一段1亚段(龙¹₁亚段)1小层海相页岩气甜点段作为研究对象, 开展系统的岩心观察、薄片鉴定、全岩-黏土X射线衍射、有机地球化学分析、扫描电镜观察、N₂低温吸附、CH₄等温吸附以及主、微量元素分析实验, 系统对比研究了海-陆过渡相与海相页岩气甜点段的特征与形成机理。研究结果表明: ①海相页岩气甜点段稳定分布与发育; 海-陆过渡相页岩气甜点段横向不连续, 纵向多层段发育。②海-陆过渡相页岩气甜点段总有机碳含量(TOC)高, 处于中-高成熟阶段, 有机质以Ⅱ₂-Ⅲ型为主; 海相页岩气甜点段TOC较高, 处于高-过成熟阶段, 有机质以Ⅰ-Ⅱ₁型为主。③海-陆过渡相页岩气甜点段矿物组成以黏土矿物为主, 黏土矿物孔隙以介孔和宏孔为主, 它们控制游离气的赋存; 有机质以发育微孔为主, 它是比表面积的主要贡献者, 控制吸附气的赋存。海相页岩气甜点段以石英矿物为主; 页岩有机质同时发育微孔和介孔, 它们是页岩气赋存的主要储存空间。④海-陆过渡相页岩气甜点段以吸附气为主, 平均占66.06%; 海相页岩气以游离气为主, 吸附气占11.15%~43.75%。⑤海-陆过渡相和海相页岩气甜点段有机质富集均受到古气候、古环境以及地质事件控制, 但陆源碎屑输入对海-陆过渡相页岩气甜点段有机质富集具有重要控制作用。⑥鄂尔多斯盆地海-陆过渡相页岩气单井最高产量达到7.9×10⁴ m³/d, 研究和勘探表明海-陆过渡相页岩气具有良好的勘探前景。

关键词: 差异性; 海-陆过渡相; 海相; 甜点段; 页岩气; 四川盆地; 鄂尔多斯盆地

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Different characteristics and formation mechanisms of transitional and marine shale gas sweet spots

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Abstract: To provide guidance on the exploration and production of transitional shale gas, we investigate the sweet spots of transitional shale gas in the 3rd submember of the 2nd member of the Shanxi Formation (the Shan 2³ submember) in the Daning-Jixian block along the eastern margin of the Ordos Basin and those of marine shale gas in the 1st sublayer of the 1st submember of the 1st member of the Longmaxi Formation (the Long 1¹ submember) in the southern Sichuan Basin. A combination of core and thin section observations, whole-rock and clay mineralogy by X-ray diffraction (XRD), organic geochemical analysis, scanning electron microscopy (SEM), N₂ adsorption, CH₄ isothermal adsorption, and major and trace element analyses, is applied to conduct a systematic comparative study on the characteristics and formation mechanisms of these sweet spots. The results indicate that the sweet spots of marine shale gas exhibit stable distributions

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and consistent development, while those of transitional shale gas show lateral discontinuities and occur across multiple layers vertically. The sweet spots of transitional shale gas feature high total organic carbon (TOC) content, medium to high maturity, and gas-prone organic matter of kerogen type II₂-III. In contrast, the sweet spots of marine shale gas are characterized by relatively high TOC content, high to over maturity, and oil-prone organic matter of kerogen type I-II₁. In the sweet spots of transitional shale gas, clay minerals are prevalent, where mesopores and macropores take a larger portion governing the occurrence of free gas. Organic matter in these sweet spots principally exhibits micropores, which contribute significantly to the specific surface area and determine the occurrence of adsorbed gas. In contrast, the sweet spots of marine shale gas display a dominance of quartz minerals. Their organic matter contains both micropores and mesopores, which serve as primary storage spaces for shale gas. The sweet spots of transitional shale gas predominantly exhibit adsorbed gas (average proportion: 66.06%), while those of marine shale gas show predominant free gas, with adsorbed gas accounting for merely 11.15%~43.75%. The organic matter enrichment in both types of sweet spots is governed by paleoclimate, paleoenvironment, and geologic events. Moreover, terrigenous debris input also plays a significant role in the formation of transitional shale gas sweet spots. The maximum single-well production of transitional shale gas in the Ordos Basin has been determined at up to 79,000 m³/d, demonstrating promising prospects for exploring transitional shale gas in the basin.

Key words: differentiation, transitional facies, marine facies, sweet spot, shale gas, Sichuan Basin, Ordos Basin

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中国拥有丰富的页岩气资源,据2012年资源潜力调查与评价,中国页岩气地质资源量高达134.42×10¹² m³,其中海相和海-陆过渡相页岩气资源量分别占总地质资源量的43.95%和29.82%^[1]。统计数据显示,在中国已探明的常规天然气储量中,超过半数储量气源岩为海-陆过渡相页岩^[2]。中国海相页岩沉积环境稳定,页岩气已实现商业化开发,且开发效果显著,先后建成涪陵焦石坝、川南长宁、威远和昭通4个页岩气田,2023年页岩气年产量达250×10¹² m³^[3],较2013年产量增长了248×10¹² m³。海-陆过渡相页岩气最近几年勘探投入在加大并已见到良好苗头,钻探与发现主要集中在鄂尔多斯盆地。其中,鄂页1井在太原组日产气量达到了1.95×10⁴ m³,5口水平井在山西组进行压裂后,测试产量(1.18~7.90)×10⁴ m³/d,大宁一吉县地区5口直井在山西组压裂后均获工业气流,最高无阻流量达到2.30×10⁴ m³/d。但由于海-陆过渡相沉积环境变化快,纵向岩性变化复杂,导致海相与海-陆过渡相页岩气在甜点段评价上存在较大差异。针对海相页岩,已有研究表明有机质孔隙是控制页岩气含气量的主要控制因素^[4],并建立了一套针对海相页岩气评价选区的指标体系^[5-6],通过笔石带对龙马溪组小层进行精细划分,厘定了威远、长宁和涪陵等区块龙马溪组页岩气发育的甜点段^[7-8];相比之下,海-陆过

渡相页岩气的勘探研究总体还处于早期阶段,多集中在沉积特征分析^[9-11]、储层表征^[12-14]、勘探潜力^[15-17]和有利层段/区优选等方面^[18-20],研究认识到海-陆过渡相页岩气发育的有利相带为潟湖相与海湾相,储集空间以矿物转化孔、缝为主,有机质孔隙欠发育,二叠系山西组2段3亚段(山2³亚段)是山西组页岩气发育的有利层段,在鄂东缘优选出3个有利区。

当前页岩气的勘探开发,主要针对甜点段,海相龙马溪组页岩气在研究和勘探开发上都积累了丰富的经验。对比海相龙马溪组页岩气甜点段,能充分认识海-陆过渡相与海相之间的差异,更好地理解不同沉积环境下页岩气的生成和储集机制,揭示不同矿物组分对页岩气开发的影响,优选更为合理的资源评价参数体系,有助于建立更为全面的地质模型。另外,四川盆地与鄂尔多斯盆地类型不同,二者页岩气发育特征的对比可以为后期同类型盆地的页岩气研究提供参考。鉴于鄂尔多斯盆地拥有丰富的天然气资源,并已在致密砂岩气^[21]、煤层气^[22]和灰岩气^[23]领域取得重要突破,进一步强化页岩气的研究将有助于推动该领域的突破性进展,进而实现盆地天然气资源的全面开发。前期虽对海相与海-陆过渡相页岩储层差异性进行了对比^[24],但并未聚焦于甜点段的对比,且当前海相与海-陆过渡相页岩甜点段差异特征以及甜点段

形成机理差异也并未进行深入研究。为此,本文以鄂东缘山西组页岩气发育甜点段与川南龙马溪组页岩气甜点段为例,开展对比分析,从而为后期海-陆过渡相页岩气的规模开发提供理论基础。

1 地质背景与实验方法

1.1 地质概况与样品来源

鄂尔多斯盆地东缘在加里东时期遭受剥蚀,至晚石炭世再次接受沉积,本溪组沉积厚度仅为几十米,早二叠世太原组沉积期形成统一的滨浅海沉积,早二叠世山西组沉积期,海水向东南方向退出,从而进入海-陆过渡相和陆相沉积时期。大宁-吉县区块位于鄂尔多斯盆地东南缘,下二叠统山西组的山²亚段为典型的海-陆过渡相沉积,是当前海-陆过渡相页岩气勘探的重点层系^[12,16-17]。山²亚段根据电性和元素特征从

上至下又可以分为3个小层(图1a,c)。

川南地区志留系龙马溪组沉积于赫兰特冰期之后的早志留世鲁丹-埃隆期^[25]。龙马溪组与下伏的观音桥段灰岩呈整合接触。龙马溪组自下而上可以分为龙一段和龙二段,龙一段又分为龙一¹亚段和龙一²亚段,龙一¹亚段从下至上又分为4个小层(图1b,d)。按照陈旭院士^[26]提出的奥陶纪-志留纪笔石划分标准,威远地区龙一¹亚段对应LM1—LM9笔石带。本次研究主要针对川南地区龙马溪组的甜点段和大宁-吉县区块山²亚段的甜点段进行采样和开展对比分析,探讨其甜点段的特征与形成机理的差异性。

1.2 实验方法

选取威远地区威202井、威204H10-2井、自201井和泸205井龙马溪组样品93块,大宁-吉县区块山²亚段样品37块,开展总有机碳含量(TOC)测试、全岩-黏土矿物X射线衍射分析、扫描电镜观察、显微组分识

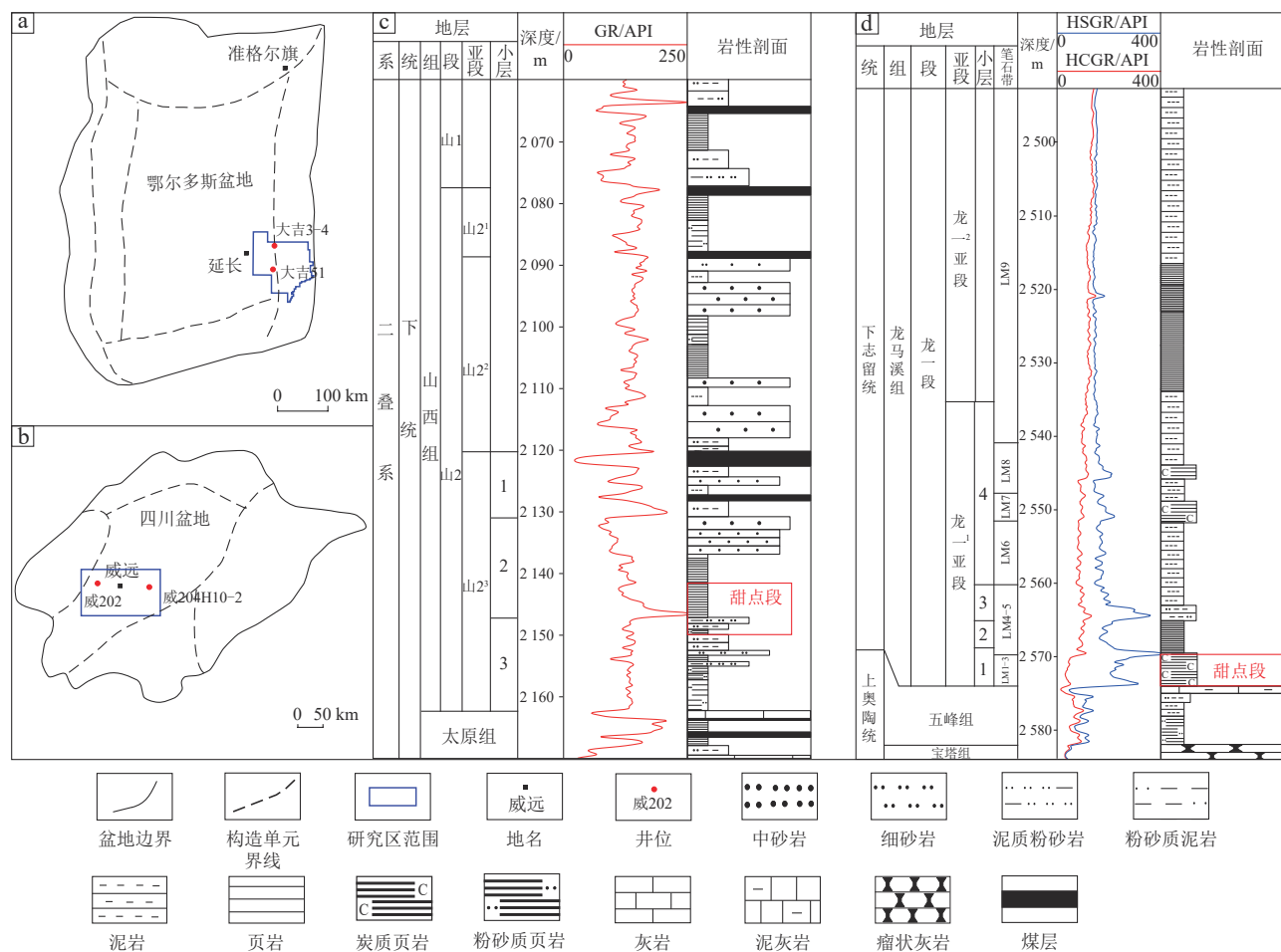


图1 鄂尔多斯盆地东缘山西组(a,c)与四川盆地南部威远页岩气田龙马溪组(b,d)综合柱状图

Fig. 1 Composite stratigraphic columns of the Shanxi Formation along the eastern margin of the Ordos Basin (a, c) and the Longmaxi Formation in the Wei yuan shale gas field of the southern Sichuan Basin (b, d)

别(其中识别标准主要参考文献[27-30])、 N_2 低温吸附、 CH_4 等温吸附以及主、微量元素分析,为系统分析海-陆过渡相和海相页岩甜点段特征提供基础数据。

2 甜点段定义

关于页岩气“甜点段”^[6,31-32],中国目前尚未有明确定义,一般是将TOC高、孔隙度高、含气量高、脆性矿物含量高和厚度适中的页岩层作为勘探开发有利层段,在实际生产中,页岩气甜点段是当前页岩气勘探开发的目标层段,具有高产、稳产能力强的特征。

针对威远地区龙一¹亚段1小层内部上、中、下储层静态参数(TOC、孔隙度、脆性矿物含量和含气量)进行了对比分析,综合评价认为储层的中、下品质优于上部:在TOC方面,1小层下部(6.70%)>中部(5.00%)>上部(3.30%);孔隙度上,1小层下部(7.8%)>中部(7.0%)>上部(5.5%);含气量上,下部(6.2 m³/t)>中部(5.6 m³/t)>上部(5.1 m³/t);脆性矿物含量上,1小层下部(89.0%)>中部(68.1%)>上部(46.5%)(图2)。

将具有“四高”特点的层段定义为威远地区龙马溪组页岩气的甜点段,即TOC>4.00%、孔隙度>6.0%、含气量>5.0 m³/t以及脆性矿物含量>60.0%。威远地区的甜点段主要分布于龙马溪组底部1小层中、下部,总体对应LM1—LM3笔石带。对生产测井资料的分析也印证了上述结论^[25],在各层的产量贡献中,1小层具有最高的产气量,是上覆2小层的2~5倍,3小层的10倍。甜点段的钻遇长度与首年日产气量呈明显正相关性,也证实了甜点段标准的准确

性和实用性。

海-陆过渡相山2³亚段页岩发育程度好,TOC高,是目前海-陆过渡相页岩气勘探的重点层段。根据实钻情况、生产效果和实验测试数据分析,将TOC>4.00%,孔隙度>2.0%,含气量>2.0 m³/t和脆性矿物含量>50.0%作为甜点段划分标准,在大吉51井区,山2³亚段的海-陆过渡相页岩气甜点段分布在3小层,在大吉3-4井区页岩气甜点段分布在2小层。该甜点段实钻也取得了突破,直井产量(1.80~2.30)×10⁴ m³/d,水平井产量(1.18~7.90)×10⁴ m³/d。在下文中,所列举的各种参数均来自海相龙一¹亚段页岩甜点段和海-陆过渡相山2³亚段页岩甜点段。

3 甜点段差异对比

3.1 连续性差异

沉积环境的差异性导致海相和海-陆过渡相页岩气甜点段在横向分布连续性上存在差异。海相龙马溪组页岩形成于稳定的深水陆棚环境,因此横向上连续性好,厚度稳定(图3)。在威远地区受到乐山-龙女寺古隆起东南斜坡水下高地影响,龙马溪组在某些区域缺失LM1—LM3笔石带,相反,在水下平原和水下洼地,这些笔石带发育完整,厚度也相对稳定(3.90~4.50 m),而水下斜坡区域,笔石带略薄(1.00~2.98 m)。总体上,甜点段呈现从东南向西北逐渐尖灭的趋势。在威210井—自204井龙马溪组8个笔石带连井剖面上,威203井、威204H10-2井、威202井和自201井笔石带发育完整,威210井LM1—LM3笔石带缺失,显

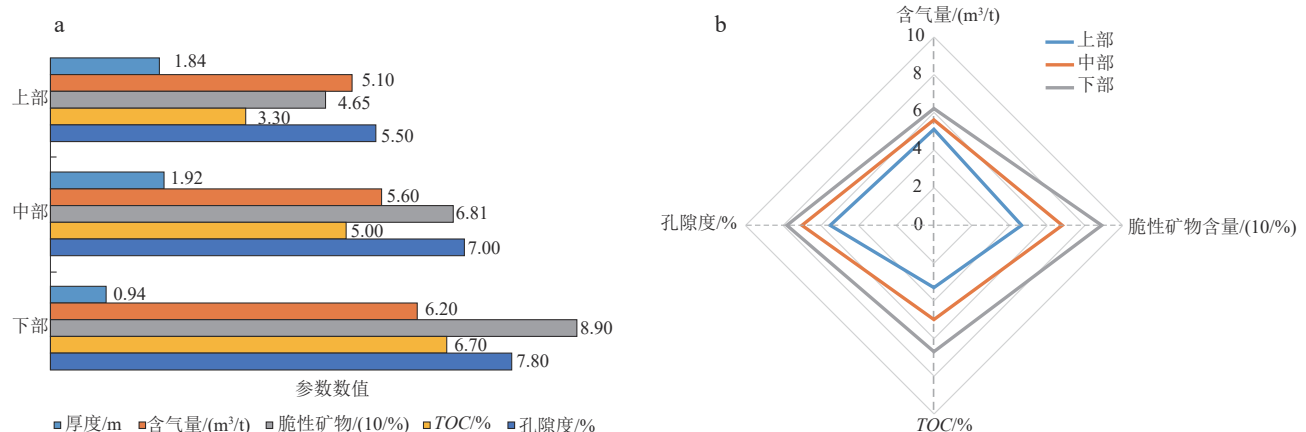


图2 四川盆地南部威远地区龙一¹亚段1小层上、中、下层段对应储层参数特征

Fig. 2 Reservoir parameters' comparison of the upper, middle, and lower sections of the 1st sublayer of the Long 1st submember in the Weiyuan area

a. 储层关键参数条形对比图; b. 储层关键参数雷达图

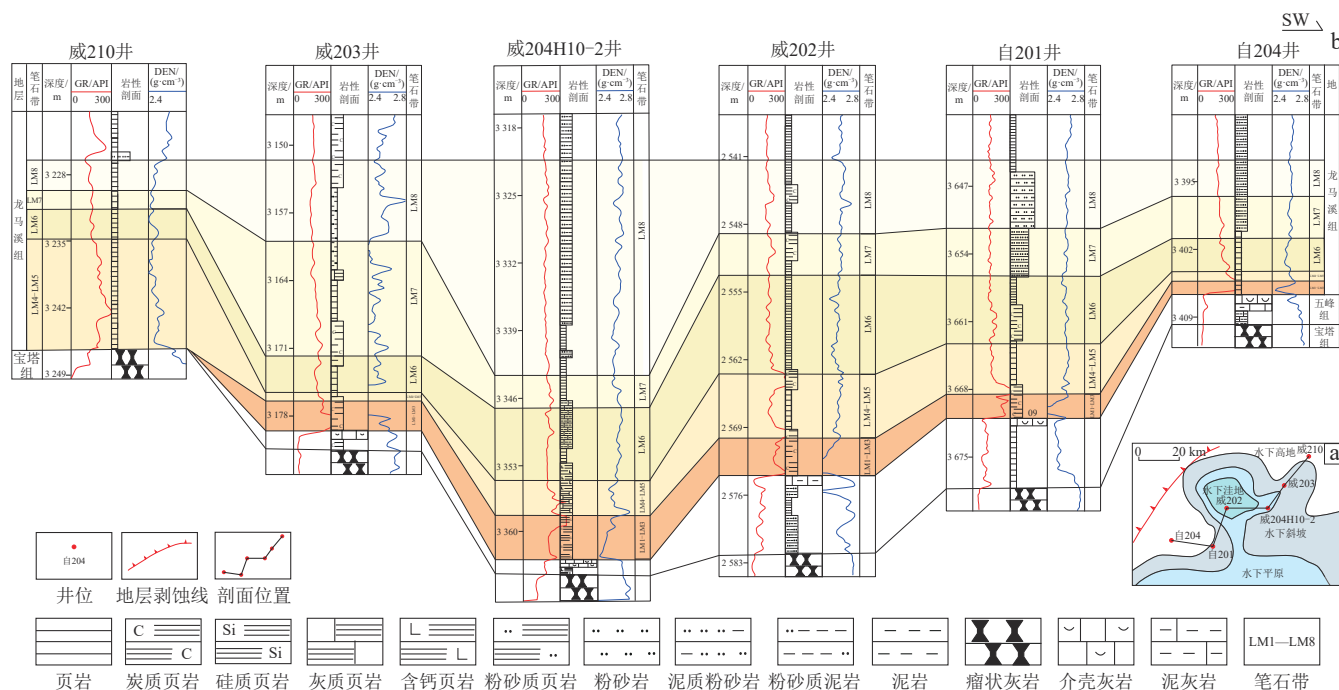


图3 四川盆地南部威远页岩气田龙马溪组甜点段横向分布特征(剖面位置图、威210笔石带均引自文献[33])

Fig. 3 Lateral distribution of the sweet spots in the Longmaxi Formation, Weiuan shale gas field (the section locations and the graptolite zones in well Wei 210 quoted from reference [33])

a. 威远地区龙马溪组不同古地貌单元平面分布; b. 威远地区龙马溪组甜点段横向展布连井剖面

示出地层逐渐上超尖灭的现象。马新华等^[25]研究也表明甜点段厚度受古地形影响,例如威202井和威204井甜点段平均厚度为3.8 m,而威205井由于受到水下古隆起的影响导致甜点段缺失,威201井由于靠近剥蚀边界而缺少龙马溪组甜点段。海-陆过渡相由

于沉积环境的快速变化和物源供给的显著影响,导致页岩甜点段发育位置并不稳定。例如在大宁—吉县区块,南部的大吉51井区山₂³亚段页岩气甜点段分布在底部3小层,厚度为4.5 m,而北部大吉3-4井区则主要发育在2小层,厚度为5.6 m(图4)。大吉51井甜点

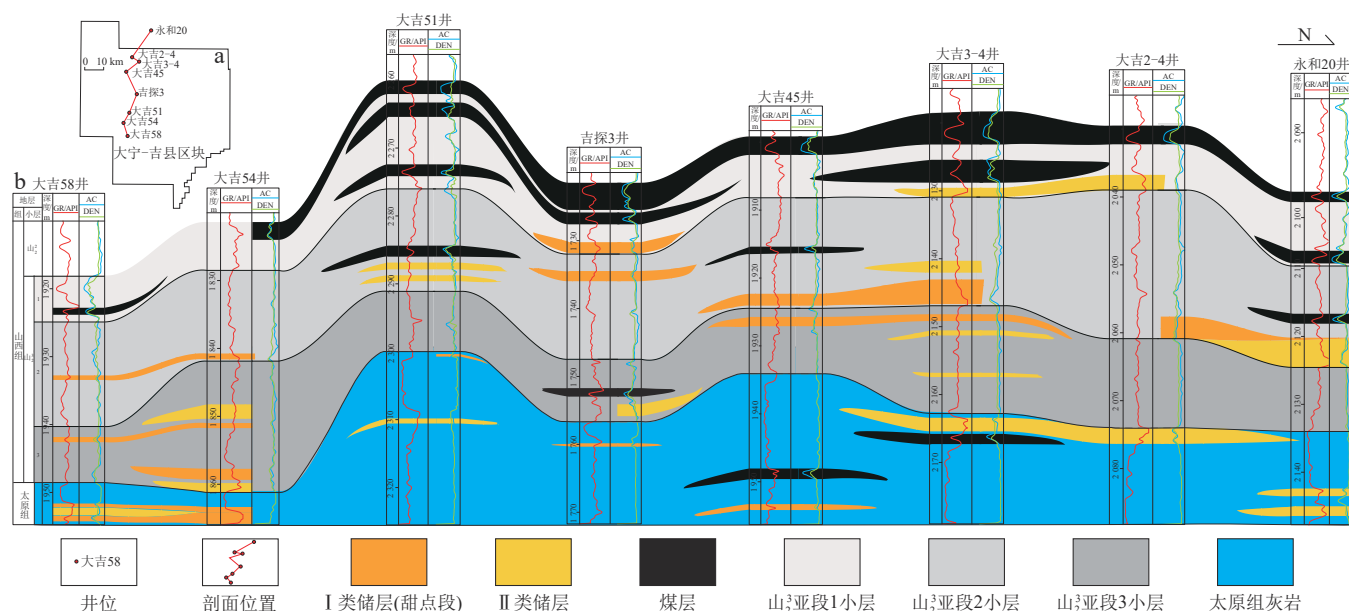


图4 鄂尔多斯盆地东缘大宁—吉县区块山₂³亚段优质储层横向与纵向分布特征

Fig. 4 Lateral and vertical distributions of high-quality reservoirs in the Shan 2³ submember, Daning-Jixian block

a. 大宁—吉县区块范围与剖面位置; b. 大宁—吉县区块南北向连井剖面特征

段富含海百合、腕足和骨针等海洋生物碎片^[10,19]。同样,在大吉3-4井的2小层岩石薄片中也观察到大量海相生物化石碎片,可鉴定的主要包括海绵骨针和介形虫(图5)。此外,由南向北的剖面上,除了2小层和3小层外,1小层也局部发育甜点段,进一步证实了海-陆过渡相甜点段纵向位置分布的不稳定性。

3.2 有机地球化学特征差异

山西组页岩的形成受到海水和陆源碎屑的共同影响,且形成时代晚,导致其在 TOC 、有机质类型以及成熟度上均与威远龙马溪组页岩存在显著差异。龙马溪组 TOC 分布在3.22%~7.83%,平均值为4.82%,镜质体反射率(R_o)介于2.76%~2.82%^[24],干酪根碳同位素值($\delta^{13}C$)分布在-30.37‰~-29.35‰,平均值为-29.86‰,有机质类型为I-II₁型(划分标准参照文献[34])。海-陆过渡相大吉51井 TOC 分布在4.53%~11.68%,平均值为7.98%;大吉3-4井 TOC 分布在4.19%~34.70%,平均值为15.64%(图6)。山2³亚段页岩气甜点段有机质类型主要为II₂型,并含有一定量III型。山2³亚段页岩气甜点段 R_o 分布在2.02%~2.61%,平均值为2.32%。对比之下,海相页岩 TOC 分布均匀,最高值较小(7.70%),而海-陆过渡相页岩

TOC 跨度大,最高值明显高于海相页岩(大吉3-4井34.70%,大吉51井11.68%),可能是由于海-陆过渡相页岩有一定量陆源碎屑有机质的供应,在大吉51井和大吉3-4井以及柳林成家庄剖面均见到丰富的植物碎片化石(图7),是陆源碎屑有机质输入的直接证据。海相页岩成熟度明显高于海-陆过渡相页岩,甚至达到了石墨化阶段^[35]。由于壳质组和镜质组的生烃能力较腐泥组差,因此在相似的热演化程度以及相同的 TOC 下,海-陆过渡相页岩的生烃潜力弱于海相页岩。

3.3 储层特征差异

3.3.1 矿物组成差异

海相页岩矿物组成以脆性矿物为主,黏土矿物较少。威远地区龙马溪组矿物组成以石英为主,含量在42.2%~91.8%,平均值为77.5%^[36],黏土矿物含量较低,平均值仅为3.0%,其中伊利石占主体(含量>80.0%)^[24]。海-陆过渡相页岩矿物组成与海相页岩差距较大,主要体现在石英含量低(10.7%~61.8%,平均值为38.1%),黏土矿物含量高(26.8%~61.1%,平均值50.2%),且黏土矿物中高岭石、伊/蒙混层与伊利石

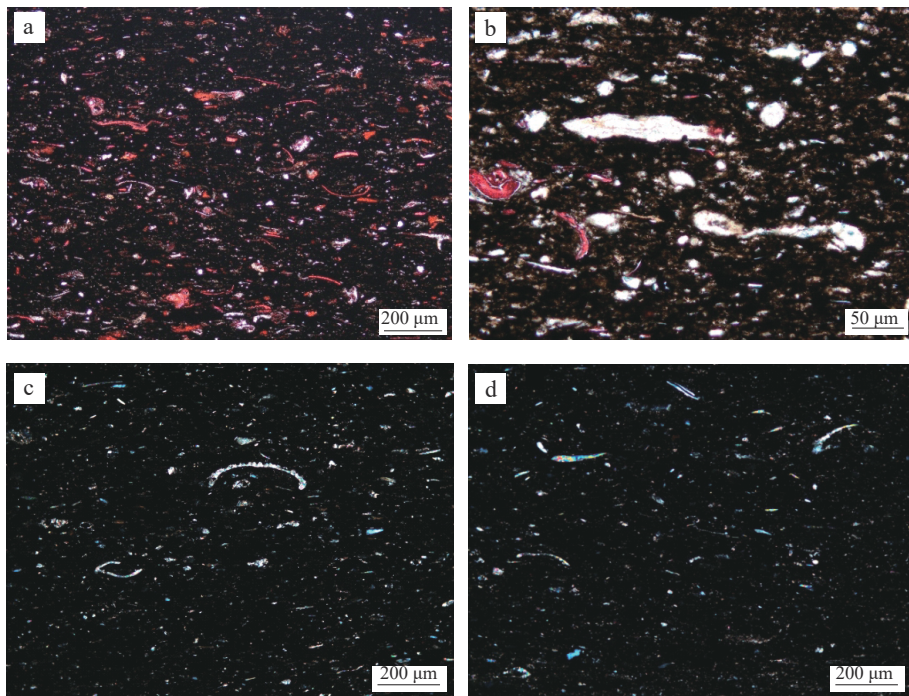


图5 鄂尔多斯盆地东缘大宁—吉县区块山2³亚段生物碎屑岩石薄片显微照片

Fig. 5 Bioclasts detected in thin sections from the Shan 2³ submember in wells Daji 51 (a, b) and Daji 3-4 (c, d), Daning-Jixian block along the eastern margin of the Ordos Basin

- a. 大吉51井,埋深2 294.74 m,正交偏光,介形虫和腕足生物碎片;b. 大吉51井,埋深2 294.20 m,正交偏光,介形虫和腕足生物碎片,生屑重结晶;
c. 大吉3-4井,埋深2 145.21 m,正交偏光,腕足和介形虫生物屑;d. 大吉3-4井,埋深2 144.93 m,正交偏光,有孔虫和介形虫生物屑

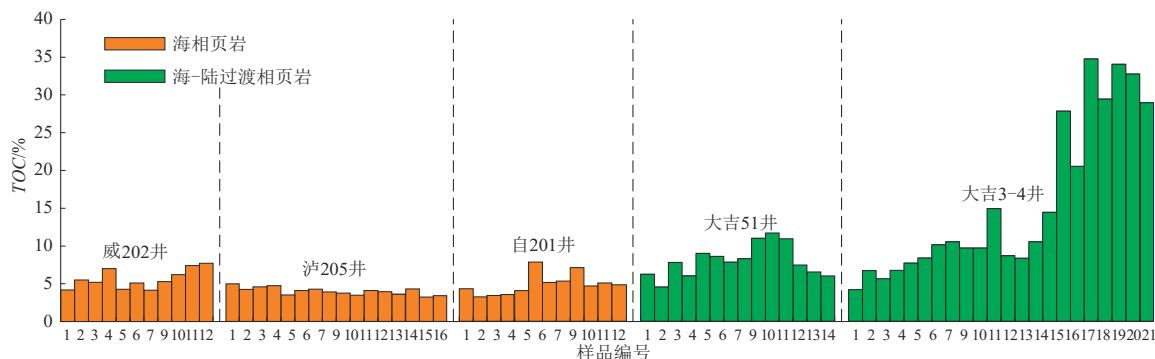


图6 四川盆地南部海相龙马溪组与鄂尔多斯盆地东缘大宁—吉县区块海—陆过渡相山西组页岩气甜点段TOC特征对比(威202井样品部分引自文献[36],大吉51井样品引自文献[16],大吉3-4井样品引自文献[12])

Fig. 6 Comparison of TOC content in sweet spots of the marine Longmaxi Formation in the southern Sichuan Basin and the transitional Shanxi Formation in the Daning-Jixian block along the eastern margin of the Ordos Basin (partial samples of well Wei 202 from reference [36], samples of well Daji 51 from reference [16], and samples of well Daji 3-4 from reference [12])

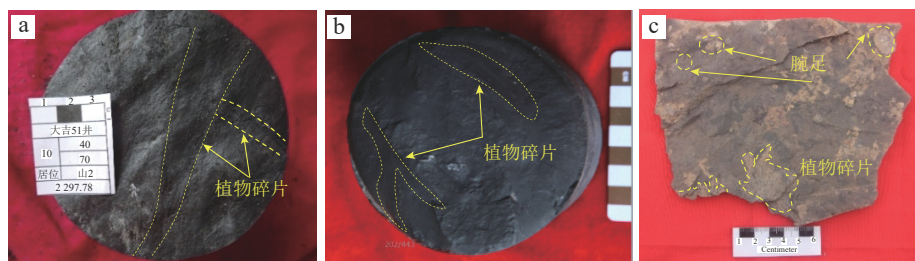


图7 鄂尔多斯盆地东缘山2³亚段植物碎片化石特征照片

Fig. 7 Characteristics of plant debris in core and outcrop samples from the Shan 2³ submember along the eastern margin of the Ordos Basin
a. 大吉51井,埋深2 297.78 m,岩心,植物碎屑呈条带状分布;b. 大吉3-4井,埋深146.00 m,岩心,植物碎片呈条带状和片状分布,植物叶片与腕足化石在层面分布;c. 柳林成家庄剖面,露头样品

含量三者相当,平均值分别为34.0%,31.0%和29.0%,并同时含有少量绿泥石(含量<10.0%)^[12]。

3.3.2 孔隙特征差异

海—陆过渡相页岩孔隙度普遍低于海相页岩。例如,大吉3-4井页岩气甜点段柱塞样的氦气孔隙度在1.59%~2.11%,平均值为1.80%^[37];大吉51井氦气孔隙度分布在0.16%~2.71%,平均值为0.94%^[18]。相比之下,海相威202井柱塞样孔隙度分布在6.40%~8.50%,平均值为7.32%^[36];威204H10-2井孔隙度分布在4.42%~8.75%,平均值为5.89%。

海相页岩有机质主要以分散状充填在矿物颗粒之间(图8a—c),部分沿层理条带状分布(图8d)。显微组分分析显示,有机质主要为藻类体、无定型组分以及少部分焦沥青,但无论是大片的形态有机质(动物碎屑如笔石、几丁虫),还是小范围分布的无定型有机质,均观察到大量近圆形的有机质孔(图8c,e,f),孔径多分布在10~250 nm,并以10~30 nm的介孔为主^[38]。不同区块的有机质面孔率略有差异,但大致在10.00%~

50.00%,平均约20.00%^[39-40]。N₂低温吸附实验结果表明,海相页岩有机质孔在孔隙系统中占主导地位,有机质同时发育微孔和介孔,以介孔为主,约为48.58%,其次为微孔,约23.62%^[41];孔隙比表面积微孔占绝对优势,约占76.87%,其次是介孔,约占23.05%^[42]。

海—陆过渡相页岩中有机质主要以条带状和块状分布在矿物基质中。镜质组为主要显微组成,边缘平直(图8g),其次为固体沥青,无定型有机质罕见,丝质体局部发育,具有胞腔结构,但基本都被矿物充填(图8h—j)。扫描电镜观察显示,海—陆过渡相有机质孔隙发育程度远不如海相页岩,丝质体中未见孔隙,镜质体在高倍放大下可见少量分散、不规则发育的孔隙,孔径分布在10~200 nm,主峰孔径为32 nm(图8k,l)。Zhang等^[13]通过干酪根分析发现,有机质孔隙比表面积占样品总孔隙比表面积的73.2%,而孔体积仅占样品总孔体积的33.5%。TOC与微孔孔隙比表面积呈强正相关,相关系数(R^2)达到0.874 0,TOC与介孔孔隙比表面积无相关性,TOC与总孔隙比表面积呈强正相关性,相关系数(R^2)为0.829 3(图9),反映海—陆过

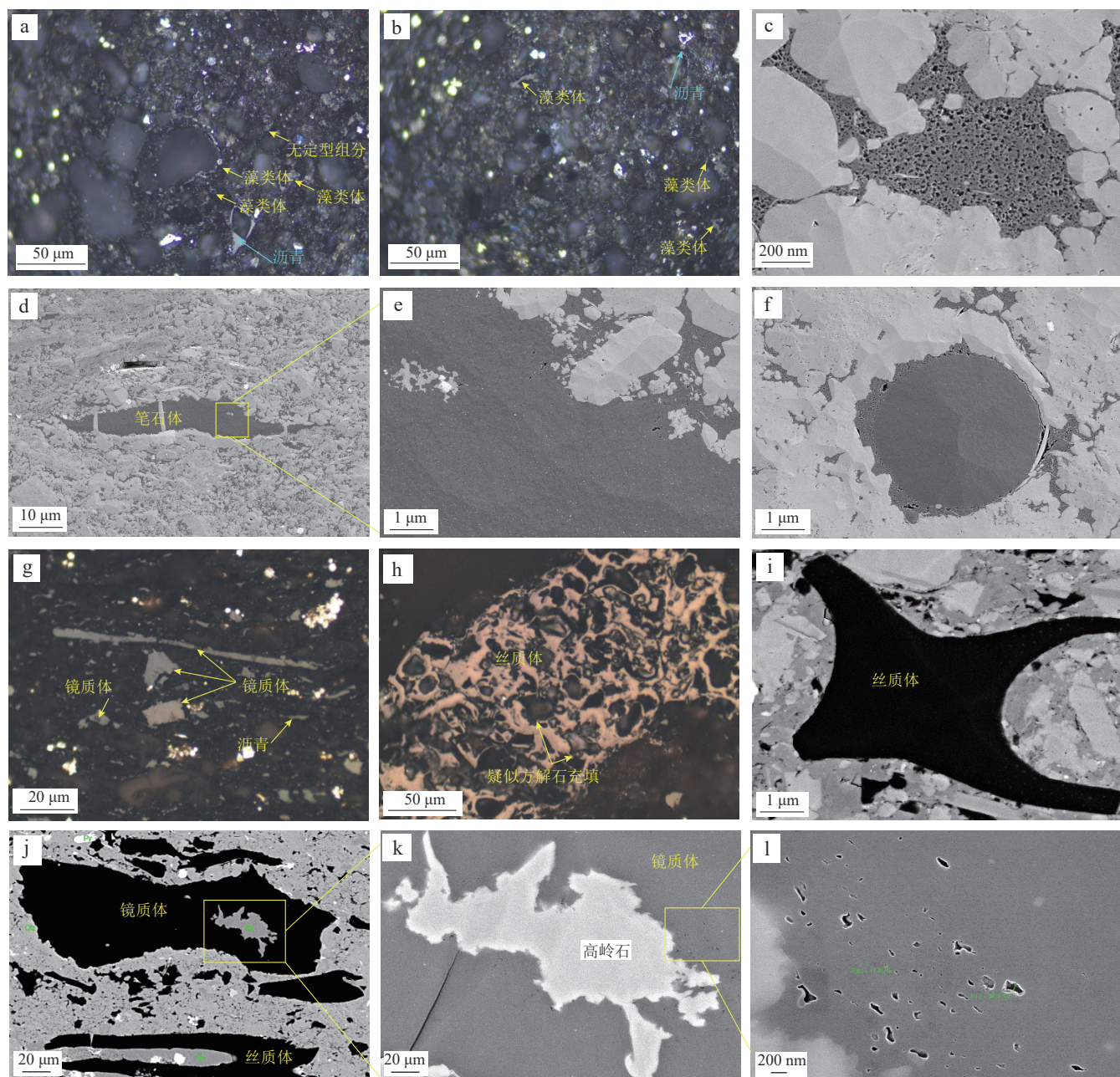


图8 四川盆地南部海相龙马溪组与鄂尔多斯盆地东缘大宁—吉县区块海-陆过渡相山西组页岩气甜点段有机质分布与孔隙发育特征显微照片

Fig. 8 Characteristics of organic matter distribution and pore development in shale gas sweet spots of the marine Longmaxi Formation in the southern Sichuan Basin and the transitional Shanxi Formation in the Daning-Jixian block along the eastern margin of the Ordos Basin

a—f. 海相有机质及其孔隙特征: a, b. 泸205井, 埋深4 030.36 m, $TOC=5.00\%$, 藻类体和无定型组分分布; c. 自201井, 埋深3 669.80 m, $TOC=4.67\%$, 无定型分布的有机质孔隙发育; d. 自201井, 埋深3 670.80 m, $TOC=7.83\%$, 顺层分布的笔石体; e. 为d图中笔石体局部放大, 孔隙发育明显; f. 自201井, 埋深3 670.55 m, $TOC=5.15\%$, 生物壳体被有机质充填, 蜂窝状孔隙发育; g—l. 海-陆过渡相有机质及其孔隙特征: g. 大吉51井, 埋深2 295.84 m, $TOC=5.98\%$, 镜质体边缘界限清晰, 呈灰黑色, 局部见沥青体; h. 大吉51井, 埋深2 298.45 m, $TOC=9.96\%$, 丝质体, 具有胞腔结构, 被疑似方解石充填; i. 大吉51井, 埋深2 295.84 m, $TOC=5.98\%$, 丝质体, 具胞腔结构; j. 大吉51井, 埋深2 293.61 m, $TOC=8.96\%$, 镜质体与丝质体, 镜质体边缘界限清晰, 在内部发育高岭石, 丝质体的胞腔结构被矿物填充, 具有圆滑的弧形边界; k. 为j图局部放大, 高岭石周边发育有机质孔隙; l. 为k图局部放大, 有机质孔隙形态不规则, 菱角状, 孔隙大小不一, 孔径10~100 nm范围内孔隙均发育

渡相有机质主要发育微孔, 从而增大页岩比表面积。从另一个角度揭示, 镜质组虽不利于发育介孔等较大孔隙, 但是其微孔发育能力显著。

综上, 海-陆过渡相页岩有机质微孔大量发育, 介孔与宏孔主要来源于黏土矿物。而海相页岩, 介孔和微孔均发育于有机质, 其 N_2 低温吸附曲线滞后环, 根

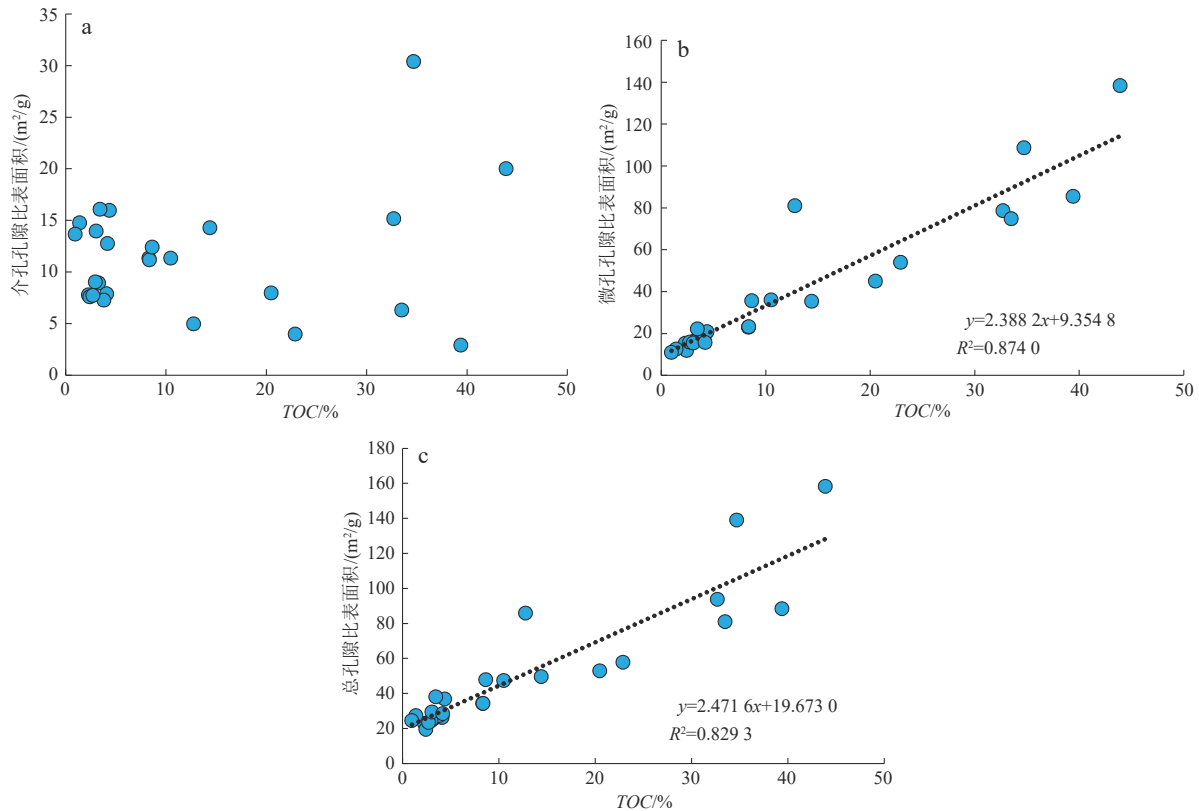


图9 鄂尔多斯盆地东缘大宁—吉县区块山²₃亚段页岩有机质对孔隙发育的控制作用

Fig. 9 Controlling effects of organic matter on pore development in the Shan ²₃ submember shale in the Daning-Jixian block along the eastern margin of the Ordos Basin

a. 有机质丰度与介孔孔隙比表面积的关系; b. 有机质丰度与微孔孔隙比表面积的关系; c. 有机质丰度与总孔隙比表面积的关系

据IUPAC分类属于H₂型,H₂滞后环的特点是具有“大肚子”形态,对应于墨水瓶型的孔隙(图10a)。海-陆过渡相页岩滞后环属于H₃型,这类滞后环的特点是在相对压力接近饱和蒸气压时吸附等温线不会出现平台,(图10b),表明孔隙以板条状为主,这也与在扫描电镜下所观察到的黏土矿物孔特征一致。

另外一个差异特征是海-陆过渡相页岩微孔发育程

度明显高于海相页岩,介孔发育程度弱于海相页岩。以大吉3-4井的海-陆过渡相页岩为例,微孔孔隙比表面积介于22.97~138.30 m²/g,平均值为49.54 m²/g,介孔孔隙比表面积介于2.91~30.40 m²/g,平均值12.60 m²/g,总孔隙比表面积分布在24.64~158.30 m²/g,平均值62.20 m²/g;微孔孔体积分布在0.00198~0.02100 cm³/g,平均值为0.00900 cm³/g,介孔孔体积分布在0.01300~

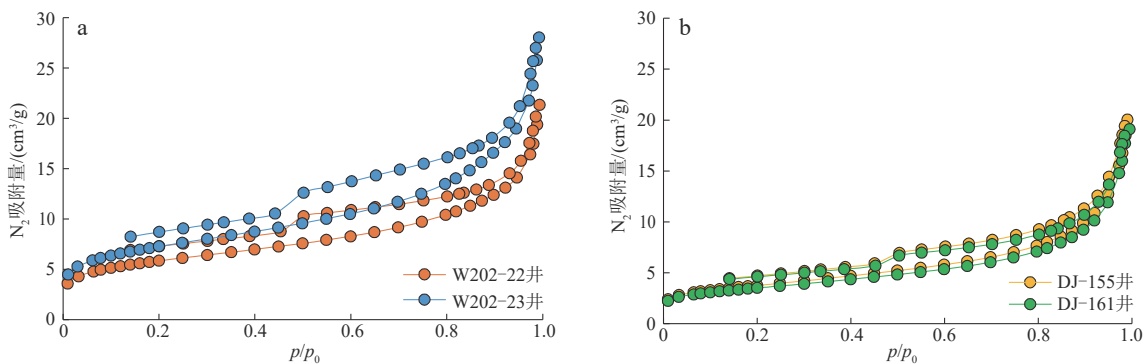


图10 四川盆地南部海相龙马溪组(a)与鄂尔多斯盆地东缘大宁—吉县区块海-陆过渡相山西组(b)页岩气甜点段样品N₂低温吸附曲线特征

Fig. 10 N₂ adsorption isotherms of sweet-spot shale samples from both the marine Longmaxi Formation in the southern Sichuan Basin (a) and the transitional Shanxi Formation in the Daning-Jixian block along the eastern margin of the Ordos Basin (b)

0.050 00 cm^3/g , 平均值 0.025 00 cm^3/g , 总孔体积分布在 0.020 00 ~ 0.066 00 cm^3/g , 平均值为 0.035 00 cm^3/g 。海相页岩以威 204H10-2 井为例, 微孔孔隙比表面积介于 5.68 ~ 13.60 m^2/g , 平均值为 9.59 m^2/g , 介孔孔隙比表面积分布在 16.91 ~ 38.31 m^2/g , 平均值为 27.14 m^2/g , 总孔隙比表面积分布在 22.75 ~ 51.92 m^2/g , 平均值为 36.73 m^2/g ; 微孔孔体积为 0.001 70 ~ 0.004 40 cm^3/g , 平均值为 0.003 00 cm^3/g , 介孔孔体积为 0.028 00 ~ 0.051 00 cm^3/g , 平均值为 0.041 00 cm^3/g , 总孔体积分布在 0.030 00 ~ 0.055 00 cm^3/g , 平均值为 0.045 00 cm^3/g (图 11)。海-陆过渡相微孔孔隙比表面积是海相页岩的 2 倍, 介孔孔隙比表面积是海相页岩的 1/2, 总孔隙比表面积是海相页岩的 2 倍; 海-陆过渡相微孔孔体积是海相页岩的 3 倍, 介孔孔体积是海相页岩的 1/2, 总

孔体积则与海相页岩相近。

3.4 页岩气赋存特征差异

页岩气以吸附态和游离态 2 种形式存在于储层中, 这 2 种相态的比例对天然气生产动态具有重要影响。威远地区龙马溪组页岩气生产曲线表现为初期高产, 随后迅速下降的特征, 单井初始最高日产气量 $(15 \sim 45) \times 10^4 \text{ m}^3$, 试采稳定日产气量 $(6 \sim 8) \times 10^4 \text{ m}^3$, 第一年产量递减率为 65 %^[43], 通过分析气体碳同位素分馏特征, 确定川东礁石坝地区页岩气吸附气占比为 13.87 % ~ 43.75 %^[44], 泸州地区页岩气吸附气占比为 11.15 % ~ 24.56 %^[45], 泸州吸附气比例较礁石坝低, 推测与该区较高的储层压力有关。

海-陆过渡相页岩具有比海相页岩更高的微孔比

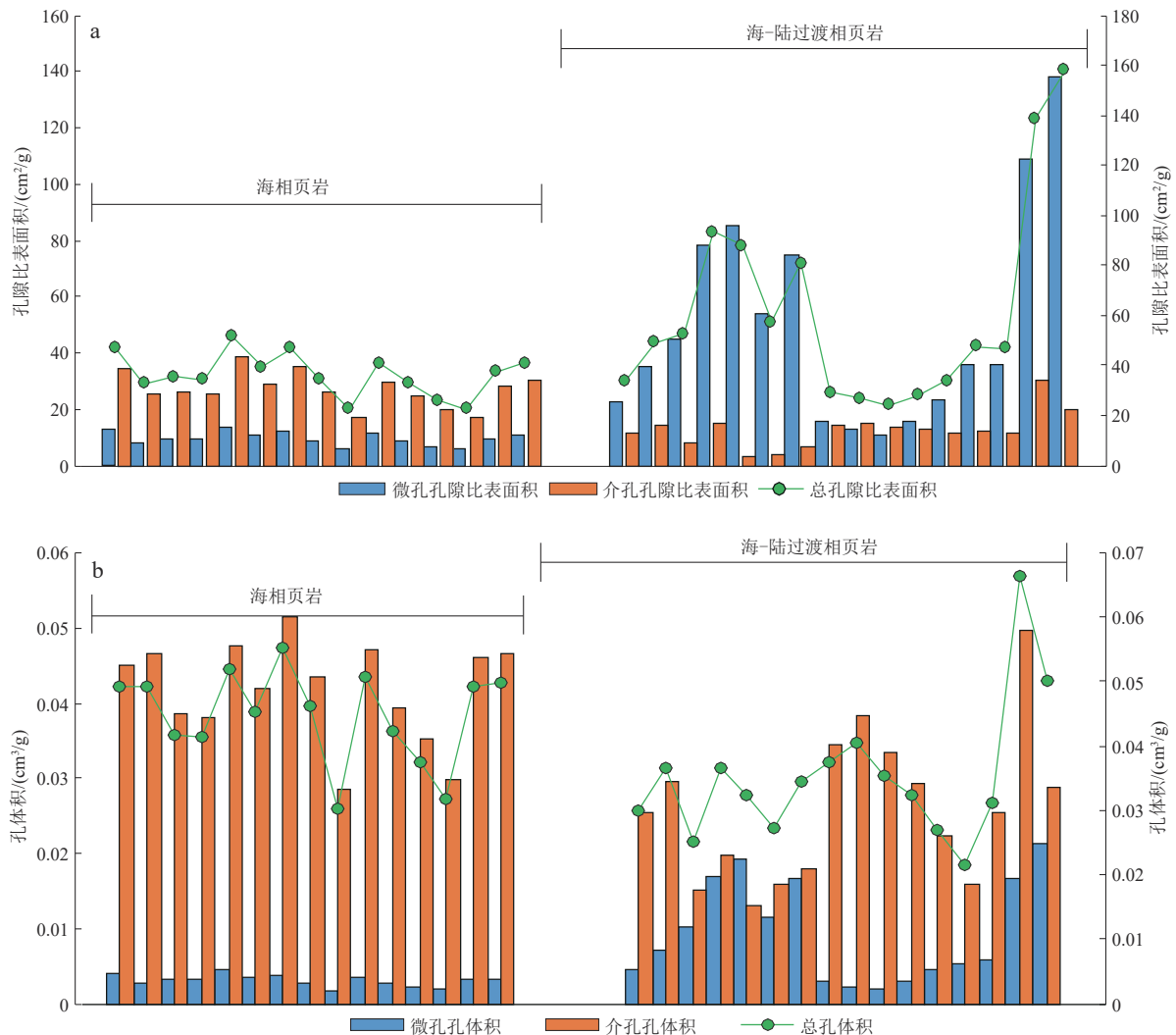


图 11 四川盆地南部海相龙马溪组与鄂尔多斯盆地东缘大宁—吉县区块海-陆过渡相山西组页岩不同尺寸孔隙比表面积(a)、孔体积(b)分布特征对比

Fig. 11 Comparison of the specific surface area (a) and volume of pores with different sizes (b) between the marine Longmaxi Formation shale in the southern Sichuan Basin and the transitional Shanxi Formation shale in the Daning-Jixian block along the eastern margin of the Ordos Basin

表面积, TOC 与兰氏体积呈强正相关性 ($R^2=0.94$)^[46], 说明微孔对吸附气控制作用明显。海相页岩兰氏体积分布在 $2.0 \sim 5.0 \text{ m}^3/\text{t}$ ^[36], 海-陆过渡相页岩兰氏体积分布在 $3.0 \sim 14.0 \text{ m}^3/\text{t}$, 平均值为 $7.5 \text{ m}^3/\text{t}$ ^[47], 远远高于海相, 进一步说明大比表面积能控制更多的吸附气含量。此外, 海-陆过渡相页岩有机质类型主要为 II_2 - III 型, 相较于 $\text{I}-\text{II}_1$ 型的海相页岩具有更强的吸附能力^[48]。海-陆过渡相页岩生产曲线表现为单井初始平均日产气量低, 分布在 $(2.0 \sim 7.9) \times 10^4 \text{ m}^3$, 产量递减缓慢的特点, 第一年产量递减率为 $20\% \sim 30\%$, 明显低于海相页岩的 65% , 说明游离气含量低。通过对大吉 54 井页岩现场解吸气量与时间的相关性分析, 采用单孔扩散模型对数据进行拟合(图 12), 单孔扩散模型计算方法参考文献^[49], 计算吸附气占比 $41.45\% \sim 87.93\%$, 平均值为 66.06% 。因此, 与海相页岩气相比, 海-陆过渡相页岩以吸附气为主。

3.5 甜点段形成机理差异

四川盆地龙马溪组页岩有机质富集机理已经得到

了广泛的研究^[50-52]。综合元素地球化学和生物地层学的分析, 学界普遍认为龙马溪组甜点段的有机质富集是古生产力和保存条件共同作用的结果(图 13a)。在沉积过程中, 诸如火山活动、冰川消融以及生物灭绝事件等均对有机质的富集产生了积极的影响^[53-54]。龙马溪组页岩成烃生物主要由疑源类和多细胞藻类构成^[7,55], 这些生物不仅具有强大的生烃能力, 还有助于有机质孔隙的形成。大规模火山喷发^[53,56-57]为硅质生物繁盛提供了必需的营养物质, 在 LM1—LM3 笔石带中, 生物硅的含量介于 $20.2\% \sim 61.5\%$, 平均值为 39.3% ^[58]。这些微晶形式的生物硅充填于陆源碎屑硅质之间, 虽然在一定程度上降低了页岩的原生孔隙度, 但与陆源碎屑硅及草莓状黄铁矿共同构成了具有高抗压实能力的骨架, 对有机质孔隙进行了有效保护。同时, 大量硅质生物的发育提高了页岩的脆性, 为水力压裂提供了有利条件。综合来看, 繁盛的水体生物和硫化缺氧环境促进了甜点段有机质富集, 有利于页岩气大量生成; 多细胞藻类的发育促进了纳米孔隙形成, 有利于页岩气的储集; 广西运动引发的海平面下降形成了封闭

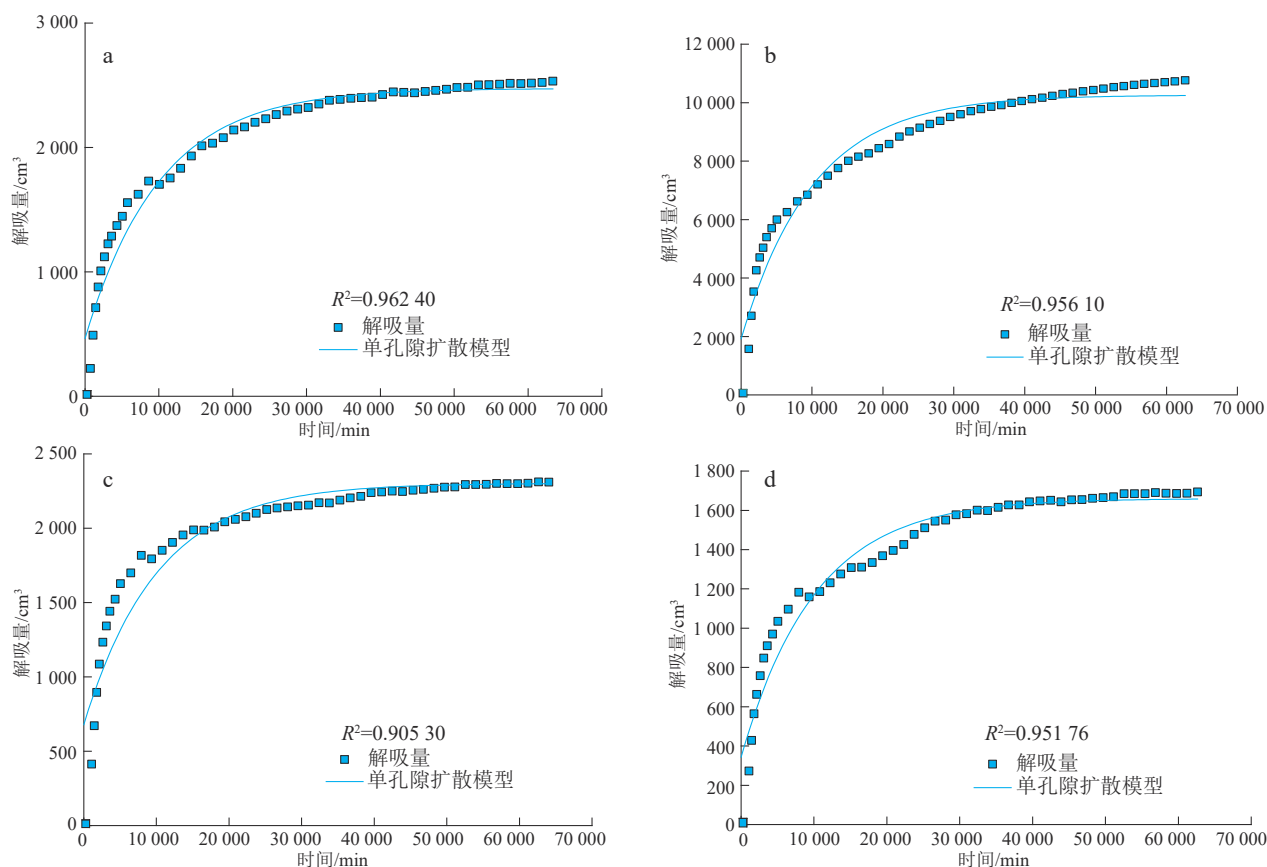


图 12 鄂尔多斯盆地东缘大宁—吉县区块大吉 54 井海-陆过渡相山西组页岩气甜点段现场解吸气量数据拟合曲线

Fig. 12 Fitted curves of on-site desorbed gas volume by uni-pore diffusion model in shale gas sweet spots of the transitional Shanxi Formation in the Daning-Jixian block along the eastern margin of the Ordos Basin

a. 埋深 1 857.2 m; b. 埋深 1 858.7 m; c. 埋深 1 858.0 m; d. 埋深 1 850.0 m

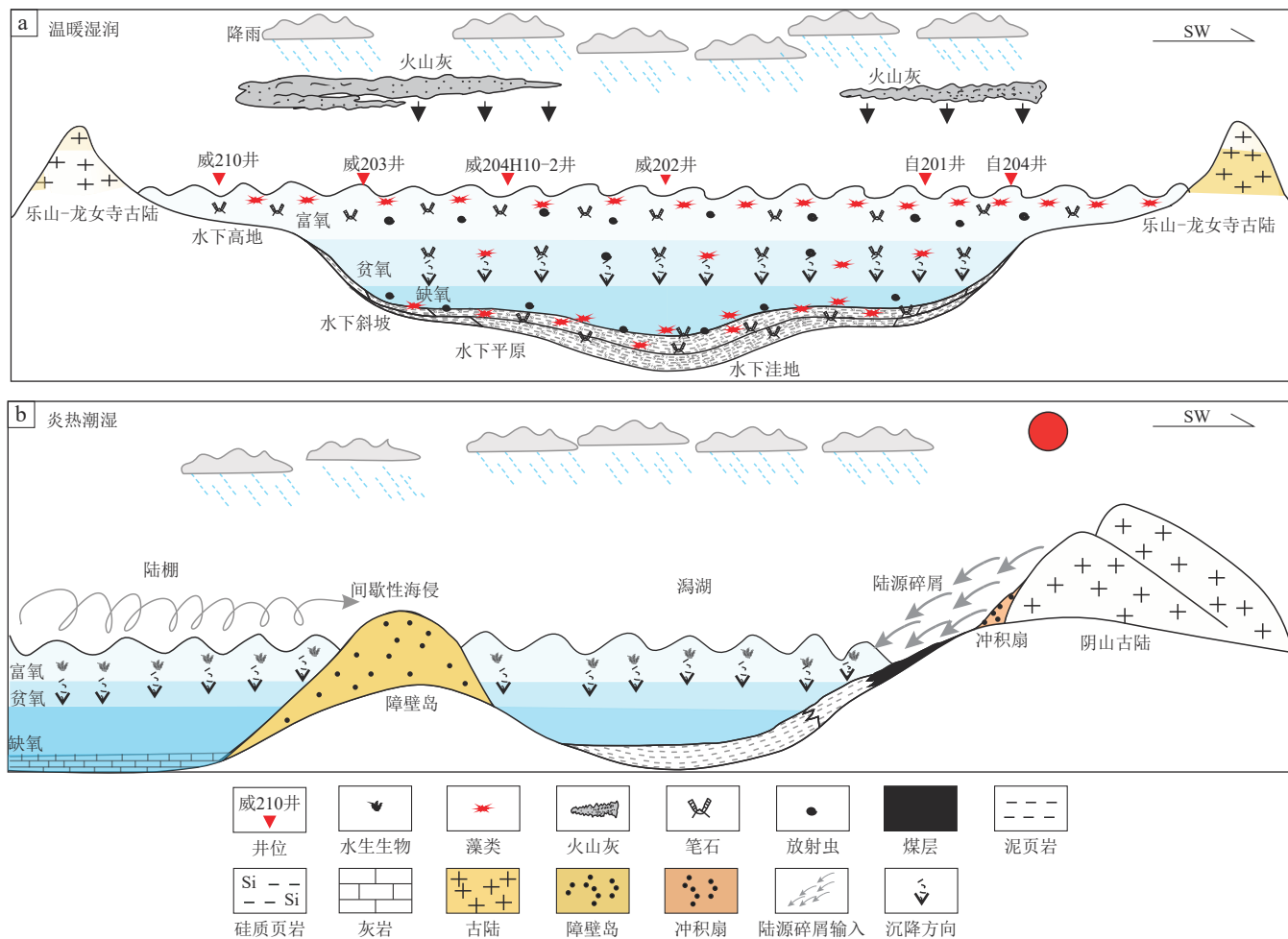


图13 四川盆地南部海相龙马溪组(a)与鄂尔多斯盆地东缘海-陆过渡相山西组(b)页岩气甜点段有机质富集模式

Fig. 13 Organic matter enrichment patterns of the shale gas sweet spots in the marine Longmaxi Formation in the southern Sichuan Basin (a) and the transitional Shanxi Formation along the eastern margin of the Ordos Basin (b)

的顶、底板,有利于页岩气的长期保存;规模火山喷发和上升洋流形成的富硅生物沉积,提高了页岩脆性和孔隙保存能力,有利于大规模体积压裂。

海-陆过渡相山西组沉积时期,华北板块以缓慢沉降为主,晚古生代冰期达到顶峰,气候变冷^[59],海平面大幅度下降,沉积界面暴露,陆地生态系统以大型维管植物为主^[60],形成高水利用率的热带雨林。这种植被结构降低了树冠的蒸腾作用,增强了地表径流^[61],风化作用增强^[59],导致大量植物碎屑和营养盐分被带入潟湖和海洋,刺激了初级生产力,从而使水体中有机碳迅速积累。海-陆过渡相有机质富集不仅受古生产力和氧化还原条件控制,还受到大量陆源碎屑输入的显著影响(图13c)。这些输入有机质主要为Ⅱ₂-Ⅲ型,分子结构中芳香族结构占比较高,生烃能力相对有限。在有机质成熟演化过程中主要发育与芳香环堆叠相关的微孔^[46],促进气体吸附。尽管有机质类型差,但由于含量高,依然能生成大量气体,并在有机质微孔中吸

附,在黏土矿物层间孔中游离赋存。山西组页岩 R_o 平均值为2.32%,接近于Ⅲ型干酪根孔隙演化的高峰阶段($R_o=2.5\%$)^[62]。由于沉积环境的频繁变化,岩性复杂多样,但潟湖相页岩质纯,在有利的上、下顶底板条件下,天然气能实现有效聚集。在大吉3-4井和大吉51井甜点段,海相生物化石的大量发现以及水体盐度的明显波动(大吉3-4井的Sr/Ba元素含量比值介于0.60~4.28,海相化石出现层段Sr/Ba>1.1),指示了间歇性海侵事件的发生,这为硅质生物的发育和脆性矿物的增加提供了条件。在大吉3-4井生物硅含量在11.30%~24.33%,平均值为17.90%,而脆性矿物含量分布在60.90%~66.90%,平均值为64.00%。综上所述,氧化还原条件、古生产力与陆源碎屑输入促进了有机质的富集;有机和无机孔隙的并行发育,促进了页岩的赋存;有利的岩性组合纵向叠置,促进了页岩的保存;间歇性的海侵带来的硅质生物提高了页岩的脆性。

4 地质意义

海-陆过渡相与海相页岩气甜点段在连续性、有机地球化学特征、储层矿物组成、孔隙特征、气体赋存特征以及甜点段形成机理上均存在显著差异(表1)。海-陆过渡相页岩相比海相页岩,具有储层连续性差,黏土矿物含量过高,储层非均质性强等不利因素,然而海-陆过渡相页岩 *TOC* 普遍高于海相页岩,热成熟度条件适中,仍然具备大规模生烃条件。鄂尔多斯盆地二叠系石盒子组大面积分布的致密砂岩气证实二叠系气源充足,虽然有机质在山西组中介孔和宏孔不如海相页岩发育,但是有机质微孔发育程度远高于海相页岩,且由于海-陆过渡相页岩以吸附气为主,这在一定程度上弱化了有机质介孔和宏孔不发育的劣势;除了有机质微孔外,山西组的无机质孔发育程度高,孔径大,有利于游离气的赋存;另外黏土矿物含量高更易于形成有机黏土矿物复合体,保护其中的有机质不易被氧化降解,有利于页岩中有机质的富集。虽然储层的连续性差,但是单层厚度可观,山₂³亚段2小层和3小层优质储层最大连续厚度能达到4.5~5.6 m,直井和水

平井均获得工业气流,最高产量能达到7.9×10⁴ m³/d,说明鄂尔多斯盆地海-陆过渡相与海相一样具有良好的页岩气勘探前景。针对海-陆过渡相页岩中黏土矿物含量高的特征,应针对性开展无水压裂实验,建立针对海-陆过渡相特有属性的评价指标体系,优选有利目标,开展攻关试验,落实资源潜力。另外,海-陆过渡相页岩与致密砂岩和煤层多形成叠置发育模式,在今后研究中应针对多种岩性组合开展综合评价,立体勘探,最大程度动用天然气有效资源。

5 结论

- 1) 海-陆过渡相页岩气甜点段具有横向不连续,纵向多层段发育的特征,而海相页岩气甜点段具有横向稳定发育,层段相对统一的特点。
- 2) 海-陆过渡相页岩气甜点段 *TOC* 异常高,热演化成熟度适中,有机质类型以Ⅱ₂-Ⅲ型为主,海相页岩气甜点段 *TOC* 较高,集中分布,热演化成熟度更高,有机质类型以Ⅰ-Ⅱ₁型为主。
- 3) 海-陆过渡相页岩气甜点段矿物组成以黏土矿物为主,有机质主要发育微孔,是比表面积的主要贡献

表1 鄂尔多斯盆地东缘大宁—吉县区块海-陆过渡相山西组与四川盆地南部海相龙马溪组页岩气甜点段差异对比

Table 1 Comparison of shale gas sweet spots between the transitional Shanxi Formation in the Daning-Jixian block along the eastern margin of the Ordos Basin and the marine Longmaxi Formation in the southern Sichuan Basin

参数	海相页岩气	海-陆过渡相页岩气
页岩气区块	威远地区	鄂东地区
沉积盆地	四川盆地	鄂尔多斯盆地
地层时代	S ₁ l	P ₁ s
埋藏深度/m	1 500~4 000	800~2 600
厚度/m	0~7.0	2.0~7.5
连续性	发育在LM1—LM3笔石带,横向稳定分布	发育在2小层与3小层,分布不连续
<i>TOC</i> /%	3.22~7.83	4.19~34.70
干酪根类型	I—Ⅱ ₂ 型	Ⅱ ₂ —Ⅲ型
<i>R_o</i> /%	2.76~2.82	2.02~2.61
矿物组分含量/%	石英为主,42.2~91.8	黏土矿物为主,26.8~61.1
黏土矿物组分含量/%	伊利石为主,>80	高岭石、伊/蒙混层和伊利石三者相当
孔隙特征	有机质孔为主,介孔和微孔均发育	有机质微孔为主,介孔和宏孔为无机质孔隙
气体赋存特征	游离态为主,占比>50%	吸附态为主,占比>60%
含气量/(m ³ /t)	2.30~8.85 ^[63] ,平均5.60 ^[64]	0.17~4.05,平均1.54
吸附气含量/(m ³ /t)	2.00~5.00 ^[36]	3.00~14.00 ^[47]
形成机理	古生产力、氧化还原条件、火山喷发和上升洋流等综合影响	古生产力、氧化还原条件、陆源碎屑输入和间歇性海侵等综合影响
地层压力系数	1.20~1.96 ^[65]	0.95~1.05

者,控制吸附气的赋存,介孔和宏孔主要由黏土矿物提供;海相页岩气甜点段矿物组成以石英为主,有机质同时发育微孔和介孔,构成页岩气赋存的主要空间。

4) 海-陆过渡相页岩气甜点段气体的赋存以吸附气为主,占总含气量的66.06%,海相页岩气甜点段以游离气为主,吸附气占比较低,不同区块虽有差异,但均低于50%。

5) 海-陆过渡相页岩气甜点段的形成受到古生产力、氧化还原条件及陆源碎屑输入的共同影响,而海相页岩气甜点段的形成主要受古生产力、氧化还原条件以及地质事件协同控制。

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